## **Depositional Environments Lec. 10 Carbonate platforms**



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# Carbonate platform







**Figure 10.2** El Capitan and Guadalupe Peak, West Texas, USA, composed of ~400 m of massive Permian fore-reef dolomite and limestone overlying gentle slopes of basinal sandstone.



## Outline









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## Shoal rimmed platform



## Open platform



## Ramp platform



## Carbonate platform











**Figure 10.1** Sketch of the general attributes of a rimmed platform with highly differentiated facies, an open unrimmed platform where high-energy conditions prevail across the shallow seafloor creating a wide subtidal facies, and a ramp over which storms sweep into shallow waters. Most complex facies are located on the narrow inner part of the structure. Source: Adapted from James et al. (2010). Reproduced with permission of the Geological Association of Canada.



Figure 10.3 An isometric sketch of facies on a reef-rimmed platform. Source: James (1983). Reproduced with permission of the American Association of Petroleum Geologists.





**Figure 10.4** An isometric sketch of facies on an ooid or skeletal sand shoal-rimmed platform. Source: James (1983). Reproduced with permission of the American Association of Petroleum Geologists.









#### **Slope Facies**



Deposits form via a variety of sediment gravity-flow processes that result in a suite of calciturbidites and calcareous debrites interbedded with finely laminated to burrowed muds



**Figure 10.12** Cambro-Ordovician slope deposits in western Newfoundland, Canada, comprising ribbon limestones at left and a massive debrite with white platform margin blocks above. Note person (circled) for scale.

## Rimmed platform



## Rim Facies



### Platform-margin shoals of ooid or skeletal grainstone







**Figure 10.5** (a) A platform rimmed by continuous coral reef growth along the continental margin of Belize, looking south on an exceptionally calm day, with various facies named. Scale bar 50 m. (b) As for (a), rimmed by a segmented coral reef. Photographs by W. Martindale. Reproduced with permission. (c) The same barrier reef on a stormy day, where all of the wave energy is absorbed by the reef rim leaving the lagoon behind largely unaffected and tranquil.









**Figure 10.6** (a) An isometric diagram of a carbonate bank with well-developed reefs on the windward side and a poorly developed rim on the leeward side, due to saline and nutrient-rich waters formed in the lagoon by rainfall and evaporation. Source: Adapted from Hanford and Loucks (1993). Reproduced with permission of the American Association of Petroleum Geologists. (b) An aerial view of the leeward margin of Glovers Atoll, Belize with a poorly developed reef (compare to Figure 10.5a, b) that is leeward facing and swept by somewhat nutrient-rich warm and saline waters. Photograph by W. Martindale. Reproduced with permission.



## **Lagoonal Facies**

can be divided into inner and outer shelf lagoons.

Muddy sediments and scattered algae, but few corals, typically characterize the inner lagoon where *Thalassia* seagrass banks are extensive. Mangroves commonly grow along the coastal boundary of this zone. Patch reefs and skeletal sands formed from skeletons of the resident calcareous algae, and invertebrates typify the outer lagoon.

**Figure 10.7** (a) A shallow subtidal lagoon community of corals and green algae in ~1m of water, Harrington Sound, Bermuda. Image width 30 cm. (b) A shallow subtidal shelf facies in ~2m of water with sparse seagrass and echinoids, Belize. Image width 20 cm image. Photographs by W. Martindale. Reproduced with permission. (c) A shallow Paleozoic subtidal facies of brachiopods (Br) and stick-like bryozoans (By) in an Upper Ordovician limestone, Anticosti Island, Québec, Canada. Image width 15 cm.





### **Beach Facies** (Beaches are common peritidal deposits)



The deposits are characterized by gentle seaward-dipping (5–15°) large-scale planar accretion beds. The *foreshore* portion is typified by alternating coarse- and fine-grained laminations, local inverse-graded laminations, and vertical burrows formed by mollusks, worms, and crustaceans



#### SIMPLE BEACH PROFILE

Figure 10.9 A sketch illustrating the different zones on a typical carbonate beach.





Ooid sand shoals, consisting of particles generated as tidal currents sweep back and forth across the shallow subtidal and intertidal sand bodies, accumulate rapidly to form segmented barriers along the edge of the platform.

**Figure 10.10** (a) A long carbonate beach with shallow foreshore (F) and offshore grassbed (G), southern Yorke Peninsula, Australia. Foreground image width 30 m. (b) Dark beachrock on a small island in the Bahamas (people for scale in background). (c) Keystone vugs (large holes) in a Holocene beach outcrop, Bahamas. Image width 8 cm. Photograph by W. Martindale. Reproduced with permission.





 fine-grained sediment production and landward transport on warmwater rimmed platforms

Deposits are replete with microbial mats, a high abundance of low-diversity biota (particularly gastropods), storm deposits, desiccation features, dolomites, and evaporites



**Figure 10.8** A tidal creek winding its way across the muddy tidal flats to the open ocean (left) in the Turks and Caicos Islands, bordered by dark intertidal microbial mats (M) and a wide supratidal zone (S). Image foreground width 80 m.



Shelf-

edge reef

oonal lime mud

Characteristic meter-scale cyclicity in platform interior



Primary relief on surface mantled and infilled by subsequent deposition
Shelf-edge skeletal sands





- Cycles may be allocyclic, driven by rising base level and creation of accommodation space for sediment accumulation
- Typical carbonate accumulate rates are very high (> base level rise), so rapidly fill accommodation space and shallow upward





#### **Aeolinite Facies**



- Holocene and Pleistocene marginal marine carbonate sand dunes (aeolianites) occur towards the northern and southern margins of the warm-water carbonate realm where the air is dry.
- Carbonate dunes are best developed where ocean waves sweep unimpeded onshore, mobilizing shallow sands and building beaches in the intertidal zone. Sand is then transported by trade winds to build dunes.
- The deposits are shore-parallel bodies in the form of transverse ridges in which the dunes are predominantly oblique, parabolic, or barchanoid. Large-scale landwarddipping foresets, grainflows, pinstripe laminations, slump scars, animal tracks, and numerous trough cross-beds related to blowouts along windward margins characterize these deposits.

#### **Aeolinite Facies**





Cliff section of aeolianite ~8 m high along south coast illustrating steep landward-dipping beds (Bermuda) Eroded trough cross-bedded aeolianites, Castle Island. Cliff is ~8 m high.





### **Supratidal flat Facies**

- Exposure (karst) surface (Carboniferous, Nevada)
- Primary relief on surface mantled and infilled by subsequent deposition



# Open platform





- These platforms lack a raised rim and so large ocean waves and swells can easily sweep across the whole shelf or bank. They are most prominent today in coolwater heterozoan settings where the waters are too cold to allow reefs or ooid sands to develop.
- No rim, no muddy open lagoon and the platform is characterized by grainy facies throughout, including beaches and aeolianites. These coarse sediments can be ooids or biofragments with microbialites in some tropical systems.
- The large amounts of sediment that are transported off the platform generate thick prograding clinoforms of predominantly neritic origin.











<u>Carbonate Ramps</u> consistent shallow gradient from shoreline to basin (some may be distally steepened), somewhat analogous to siliciclastic shelf







## Siliciclastic Equivalents







**Figure 10.14** Cross-section of a carbonate ramp indicating the importance of the carbonate factory and location of different tempestites.



### **Outer Ramp Facies**

- Thin packstone "tempestites" in lime mudstone (or siliciclastic mud) \*
- May be normally graded or contain HCS; gutter casts/sole marks on • base sea









- Shoals (ooids or skeletal) are much more common than in clastic settings
- Large subaqueous dunes, generating large-scale crossbedding









### **Middle Ramp Facies**



 Other sedimentary structures and facies may be similar to siliciclastics (wave ripples, SCS, low-angle parallel beach laminations, etc.)



Swaley cross-stratification (SCS)



Beach deposits Low-angle beds with hardgrounds



### **Inner Ramp Facies**





sea level -



### **Inner Ramp Facies**



Typical facies: light gray or light brown limestone/dolostone with

Crinkly or wavy lamination from microbial mats



<u>Fenestrae</u>: mm-sized voids (sometimes filled by calcite spar) formed by trapped gas bubbles or beneath microbial mats









Frontispiece Desiccation cracks in peritidal dolomite, Lower Ordovician, Newfoundland, Canada, hammer scale 15 cm long.



## Tepee Structures in high intertidal/supratidal zone









### Continued cycles of desiccation disrupt tepees, forming <u>flat pebble conglomerates</u>









Figure 10.15 (a) Cross-bedded nearshore inner-ramp crinoidal limestones, Upper Ordovician, Anticosti Island, Québec, Canada. Note person (circled) for scale. (b) Finely interlaminated limestone and dolostone, lagoonal facies, Middle Ordovician Maryland, USA. Note 10 cm long pen for scale. (c) Thinly bedded lenticular tempestites of the mid ramp, person scale, Upper Ordovician, Anticosti Island, Québec, Canada. Note person for scale. (d) Thinly interbedded muddy tempestites and shales, Upper Ordovician, Anticosti Island, Québec, Canada. Scale bar: 10 cm increments.





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